An Approximate Inverse to the Extended Born Modeling Operator

Jie Hou

The Rice Inversion Project

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TRIP

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Full Waveform Inversion :

Given $d \in \mathcal{D}$, find $m \in \mathcal{M}$ so that

 $\mathcal{F}[m] \simeq d$

• $\mathcal{M} =$ model space, $\mathcal{D} =$ data space

• $\mathcal{F}: \mathcal{M} \to \mathcal{D}$ Forward Map

Least Squares formulation :

Given $d \in \mathcal{D}$, find $m \in \mathcal{M}$ to minimize

 $J_{LS} = ||\mathcal{F}[m] - d||^2 [+\text{regularizing terms}]$

Strong nonlinearity, many local minima (descent methods fail)

 $\mathcal{M} = \mathsf{physical} \ \mathsf{model} \ \mathsf{space}$

 $\bar{\mathcal{M}} = \mathsf{bigger}$ extended model space

 $\bar{\mathcal{F}}:\bar{\mathcal{M}}\rightarrow\mathcal{D}$ extended modeling operator

Extension Property:

• $\mathcal{M} \subset \bar{\mathcal{M}}$

•
$$m \in \bar{\mathcal{M}} \to \bar{\mathcal{F}}[m] = \mathcal{F}[m]$$

2D Constant Density Acoustic Wave equation:

$$\frac{1}{\mathbf{v}^2(\mathbf{x})} \frac{\partial^2 u}{\partial t^2}(\mathbf{x}, t) - \nabla^2 u(\mathbf{x}, t) = f(\mathbf{x_s}, t)$$
$$u \equiv 0, t \ll 0$$

Forward Modeling (Solve the wave equation) :

$$\mathcal{F}[\mathbf{v}] = u(\mathbf{x_r}, t; \mathbf{x_s})$$

Linearization :

$$\mathbf{v}(\mathbf{x}) = \mathbf{v}_0(\mathbf{x}) + \delta \mathbf{v}(\mathbf{x})$$

Then

$$\mathcal{F}[\mathbf{v}] \approx \mathcal{F}[\mathbf{v}_0] + \mathbf{F}[\mathbf{v}_0]\delta\mathbf{v}$$

Born Approximation

$$F[v]\delta v(\mathbf{x_r}, t; \mathbf{x_s}) = \delta u(\mathbf{x_r}, t; \mathbf{x_s})$$

Born Modeling Operator F[v]

$$\begin{pmatrix} \frac{1}{\mathbf{v}(\mathbf{x})^2} - \nabla^2 \end{pmatrix} G(\mathbf{x}, t; \mathbf{x}_s) = \delta(t)\delta(\mathbf{x} - \mathbf{x}_s); \\ \left(\frac{1}{\mathbf{v}(\mathbf{x})^2} - \nabla^2\right)\delta u(\mathbf{x}_r, t; \mathbf{x}_s) = \frac{2\delta\mathbf{v}(\mathbf{x})}{\mathbf{v}(\mathbf{x})^3}G(\mathbf{x}, t; \mathbf{x}_s)$$

where G is Green's function, the implulse response of the medium

Assumption: Single scattering at points of discontinuity of impedance in the subsurface(No multiple scattering!)

Given smooth background velocity $v(\mathbf{x})$, seismic reflection data $d(\mathbf{x_r}, t; \mathbf{x_s})$, find perturbation model $\delta v(\mathbf{x})$ to fit the data:

 $F[v]\delta v \simeq d$

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Migration is an approximate solution of this linearized inverse problem

- Migration operator (producing image) is adjoint or transpose of modeling operator(Lailly, Tarantola, Claerbout(80's)).
- Migration operator can position reflectors correctly but with possibly incorrect amplitudes and wavelets.
- True amplitude migration is (pseudo) inverse

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Born Modeling Operator

$$F[\mathbf{v}]\delta\mathbf{v}(\mathbf{x}_{\mathbf{r}}, t; \mathbf{x}_{\mathbf{s}}) = \frac{\partial^2}{\partial t^2} \int d\mathbf{x} \int d\tau \frac{2\delta\mathbf{v}(\mathbf{x})}{\mathbf{v}^3(\mathbf{x})} G(\mathbf{x}, \mathbf{t} - \tau; \mathbf{x}_{\mathbf{r}}) G(\mathbf{x}, \tau; \mathbf{x}_{\mathbf{s}})$$

Born Modeling Operator

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The adjoint of F (migration operator) is defined by

$$\int d\mathbf{x}_{\mathbf{s}} d\mathbf{x}_{\mathbf{r}} dt (F\delta v)(\mathbf{x}_{\mathbf{r}}, t; \mathbf{x}_{\mathbf{s}}) d(\mathbf{x}_{\mathbf{r}}, t; \mathbf{x}_{\mathbf{s}}) = \int d\mathbf{x} \delta v(\mathbf{x}) (F^* d)(\mathbf{x})$$

Born Modeling Operator

$$F[\mathbf{v}]\delta\mathbf{v}(\mathbf{x}_{\mathbf{r}}, t; \mathbf{x}_{\mathbf{s}}) = \frac{\partial^2}{\partial t^2} \int d\mathbf{x} \int d\tau \frac{2\delta\mathbf{v}(\mathbf{x})}{\mathbf{v}^3(\mathbf{x})} G(\mathbf{x}, \mathbf{t} - \tau; \mathbf{x}_{\mathbf{r}}) G(\mathbf{x}, \tau; \mathbf{x}_{\mathbf{s}})$$

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$$\int d\mathbf{x}_{s} d\mathbf{x}_{r} dt (F\delta v)(\mathbf{x}_{r}, t; \mathbf{x}_{s}) d(\mathbf{x}_{r}, t; \mathbf{x}_{s}) = \int d\mathbf{x} \delta v(\mathbf{x}) (F^{*} d)(\mathbf{x})$$

Integration by parts leads to

$$F^*d(\mathbf{x}) = -\frac{2}{\nu^3(\mathbf{x})} \int d\mathbf{x_s} d\mathbf{x_r} dt d\tau G(\mathbf{x}, \tau; \mathbf{x_s}) \frac{\partial^2 d(\mathbf{x_r}, t; \mathbf{x_s})}{\partial t^2} G(\mathbf{x}, t - \tau; \mathbf{x_r})$$

Subsurface offset Extension



Subsurface Extension : 2h = Difference between subsurface scatering points (subsurface offset)

Physical meaning : action at a positive distance

Extend the operator by permiting δv to also depend on (half) offset h.

Extended Born Modeling and Migration Operator

$$\bar{F}[\mathbf{v}]\delta\mathbf{v} = \frac{\partial^2}{\partial t^2} \int d\mathbf{x} d\mathbf{h} d\tau G(\mathbf{x} + \mathbf{h}, t - \tau; \mathbf{x_r}) \frac{2\delta\mathbf{v}(\mathbf{x}, \mathbf{h})}{\mathbf{v}^3(\mathbf{x})} G(\mathbf{x} - \mathbf{h}, \tau; \mathbf{x_s})$$
$$\bar{F}^* d = -\frac{2}{\mathbf{v}^3(\mathbf{x})} \int d\mathbf{x_s} d\mathbf{x_r} dt d\tau G(\mathbf{x} - \mathbf{h}, \tau; \mathbf{x_s}) G(\mathbf{x} + \mathbf{h}, t - \tau; \mathbf{x_r}) \frac{\partial^2 d(\mathbf{x_r}, t; \mathbf{x_s})}{\partial t^2}$$

Fons ten Kroode (2012) constructed the inverse of the extended Kirchhoff Operator (in asymptotic sense) :

Fons ten Kroode,2012

$$\begin{split} \tilde{\mathcal{K}} &i = \frac{1}{2\pi} \int d\mathbf{x} d\mathbf{h} d\omega e^{-i\omega t} \mathcal{G}(\mathbf{x}_{\mathbf{r}}, \mathbf{x} + \mathbf{h}, \omega) \frac{\partial i(\mathbf{x}, \mathbf{h})}{\partial z} \mathcal{G}(\mathbf{x} - \mathbf{h}, \mathbf{x}_{\mathbf{s}}, \omega) \\ \tilde{\mathcal{I}} &d = \frac{32}{\pi v^{2}(\mathbf{x})} \int d\mathbf{x}_{\mathbf{r}} d\mathbf{x}_{\mathbf{s}} d\omega (-i\omega) \frac{\partial \mathcal{G}^{*}(\mathbf{x} + \mathbf{h}, \mathbf{x}_{\mathbf{r}}, \omega)}{\partial z_{r}} d(\mathbf{x}_{\mathbf{r}}, \mathbf{x}_{\mathbf{s}}, \omega) \frac{\partial \mathcal{G}^{*}(\mathbf{x}_{\mathbf{s}}, \mathbf{x} - \mathbf{h}, \omega)}{\partial z_{s}} \end{split}$$

(http://iopscience.iop.org/0266-5611/28/11/115013)

Can we construct a similar operator to extended Born Modeling Operator?

Construction of the Inverse Operator

Asymptotic Analysis of the Normal Operator $\bar{F}[v]^*\bar{F}[v]\delta v(\mathbf{x},h)$

Extended Born Modeling Operator and its Adjoint

$$\bar{F}[v]\delta v = \frac{\partial^2}{\partial t^2} \int d\mathbf{x} d\mathbf{h} d\tau \, G(\mathbf{x} + \mathbf{h}, t - \tau; \mathbf{x_r}) \frac{2\delta v(\mathbf{x}, \mathbf{h})}{v^3(\mathbf{x})} G(\mathbf{x} - \mathbf{h}, \tau; \mathbf{x_s})$$
$$\bar{F}^*[v]d = -\frac{2}{v^3(\mathbf{x})} \int d\mathbf{x_s} d\mathbf{x_r} dt d\tau \, G(\mathbf{x} - \mathbf{h}, \tau; \mathbf{x_s}) G(\mathbf{x} + \mathbf{h}, t - \tau; \mathbf{x_r}) \frac{\partial^2 d(\mathbf{x_r}, t; \mathbf{x_s})}{\partial t^2}$$

• Step 1 High Frequency Approximation : in 2D

$$G(\mathbf{x}_{\mathbf{s}}, \mathbf{x}, t) \cong a(\mathbf{x}_{\mathbf{s}}, \mathbf{x}) S(t - \tau(\mathbf{x}_{\mathbf{s}}, \mathbf{x})), \quad S(t) = t^{-1/2} H(t)$$

$$G(\mathbf{x}, \mathbf{x}_{\mathbf{r}}, t) \cong a(\mathbf{x}, \mathbf{x}_{\mathbf{r}}) S(t - \tau(\mathbf{x}, \mathbf{x}_{\mathbf{r}})), \quad S(t) = t^{-1/2} H(t)$$

• Step 2 Principle of Stationary Phase $\left(\frac{a_s^2 a_r^2}{\sqrt{det \; Hess}}\right)$

• Step 3 Modify adjoint operator by some Velocity-independent Filters

Key point in the derivation



$$\bar{F}^{-1}[v_0]\delta d(\mathbf{x},h) = -16|k||k'|v_0(\mathbf{x})^5 \bar{F}^*[v_0] I_t^4 D_{z_s} D_{z_r} \delta d(\mathbf{x},h)$$

•
$$k = (k_x, k_z)$$
 and $k' = (k_h, k_z)$ are the wavenumbers

- I_t is the time integral
- D_{z_s}, D_{z_r} are the source and receiver depth derivative.

Apply $D_{z_s}D_{z_r}$

TRIP



Reflector

Numerical Test I



Extended Migration Result









Resimulated Data



Data Residual (= 6.34% //observed data//)



Figure: One trace (middle) comparison between the original data(blue) and resimulated data(green). The differnce is shown as the red line.



Non-extended Inversion Result

Model Residual (= 9.74% ||model||)

$$\delta \mathbf{v}(\mathbf{x}) = \sum_{h} \delta \mathbf{v}(\mathbf{x}, h)$$

One trace Comparison



Figure: One trace (middle) comparison between the reflectivity model (blue) and non-extended inversion result (green). The differnce is shown as the red line.

Wrong Background Velocity - Extended Inversion TRIP



Nonextended Inversion





Resimulated Data for Extended Inversion



Data Residual for Extended Inversion



Background Velocity : 0.9v₀ 29.3%//original data//



Background Velocity : 1.1v₀ 15.6%//original data//

Data Residual for Non-extended Inversion



Background Velocity : 0.9v₀ 132.88%//original data//



Background Velocity : 1.1v₀ 158.77%//original data//

TR. LP

Apply $D_{z_s}D_{z_r}$ -Naive Implementation



Reflector

Apply $D_{z_s}D_{z_r}$ -Free Surface Simulation



Reflector



Numerical Test II



Extended Inversion



3000



Resimulated Data

Data Difference =10.4%//observed data//

Distance (m)

2000

1000

One trace Comparison



Figure: One trace (middle) comparison between the original data(blue) and resimulated data(green). The differnce is shown as the red line.

Non-extended Inversion





Non-extended Inversion Result



Model Difference =21.3% ||model||



Figure: One trace (middle) comparison between the reflectivity model (blue) and non-extended inversion result (green). The differnce is shown as the red line.







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Reflectivity Model



Non-extended Inversion Result

One Trace comparison





One Trace comparison



TΡ

Takeaway Messages

- Migration is a kinematic solution of the linearized inverse problem
- Subsurface offset extended RTM can be modified into an asymptotic inverse to the extended Born Modeling Operator
- The new inverse operator can approximate the least sqaure extended RTM solution
- The new inverse operator can also produce non-extended inversion, which can approximate least square RTM

- More Numerical Tests
- Replace D_{z_s}, D_{z_r} with respect to one-way operator
- Extension to 3D
- Apply this operator as a preconditioner to LSM and FWI

- William Symes and Fons ten Kroode
- TRIP Members
- TRIP Sponsors
- Thank you for listening