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Partially linearized seismic inverse problem (“velocity analysis”): given observed seismic data S^{obs} , find smooth *velocity* $v \in \mathcal{E}(X)$, $X \subset \mathbf{R}^3$ oscillatory *reflectivity* $r \in \mathcal{E}'(X)$ so that

$$F[v]r \simeq S^{\text{obs}}$$

Acoustic partially linearized model: acoustic potential field u and its perturbation δu solve

$$\left(\frac{1}{v^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) u = \delta(t) \delta(\mathbf{x} - \mathbf{x}_s), \quad \left(\frac{1}{v^2} \frac{\partial^2}{\partial t^2} - \nabla^2 \right) \delta u = 2r \nabla^2 u$$

plus suitable bdry and initial conditions.

$$F[v]r = \frac{\partial \delta u}{\partial t} \Big|_Y$$

data acquisition manifold $Y = \{(\mathbf{x}_r, t; \mathbf{x}_s)\} \subset \mathbf{R}^7$, $\dim Y \leq 5$ (many idealizations here!).

$F[v] : \mathcal{E}'(X) \rightarrow \mathcal{D}'(Y)$ is a linear map (FIO of order 1), but dependence on v is quite nonlinear, so this inverse problem is nonlinear.

Agenda:

- reformulation of inverse problem via *extensions*
- “standard processing” extension and standard VA
- the surface oriented extension and standard MVA
- the Ψ DO property and why it’s important
- global failure of the Ψ DO property for the SOE
- Claerbout’s depth oriented extension has the Ψ DO property

Extension of $F[v]$: manifold \bar{X} and maps $\chi : \mathcal{E}'(X) \rightarrow \mathcal{E}'(\bar{X})$, $\bar{F}[v] : \mathcal{E}'(\bar{X}) \rightarrow \mathcal{D}'(Y)$ so that

$$\begin{array}{ccccc}
 & & \bar{F}[v] & & \\
 & & \rightarrow & \mathcal{D}'(Y) & \\
 \chi & \uparrow & & \uparrow & \text{id} \\
 & \mathcal{E}'(X) & \rightarrow & \mathcal{D}'(Y) & \\
 & & F[v] & &
 \end{array}$$

commutes.

Invertible extension: $\bar{F}[v]$ has a *right parametrix* $\bar{G}[v]$, i.e. $I - \bar{F}[v]\bar{G}[v]$ is smoothing. [The trivial extension - $\bar{X} = X, \bar{F} = F$ - is virtually never invertible.] Also χ has a *left inverse* η .

Reformulation of inverse problem: given S^{obs} , find v so that $\bar{G}[v]S^{\text{obs}} \in \mathcal{R}(\chi)$ (implicitly determines r also!).

Example 1: Standard VA extension. Treat each CMP as if it were the result of an experiment performed over a layered medium, but permit the layers to vary with midpoint.

Thus $v = v(z), r = r(z)$ for purposes of analysis, but at the end $v = v(\mathbf{x}_m, z), r = r(\mathbf{x}_m, z)$.

$$F[v]R(\mathbf{x}_m, h, t) \simeq A(\mathbf{x}_m, h, z(\mathbf{x}_m, h, t))R(\mathbf{x}_m, z(\mathbf{x}_m, h, t))$$

Here $z(\mathbf{x}_m, h, t)$ is the inverse of the 2-way travelttime

$$t(\mathbf{x}_m, h, z) = 2\tau(\mathbf{x}_m + (h, 0, z), \mathbf{x}_m)_{v=v(\mathbf{x}_m, z)}$$

computed with the layered velocity $v(\mathbf{x}_m, z)$, i.e.

$$z(\mathbf{x}_m, h, t(\mathbf{x}_m, h, z')) = z'.$$

R is (yet another version of) “reflectivity”

$$R(\mathbf{x}_m, z) = \frac{1}{2} \frac{dr}{dz}(\mathbf{x}_m, z)$$

That is, $F[v]$ is a change of variable followed by multiplication by a smooth function. NB: industry standard practice is to use vertical traveltime t_0 instead of z for depth variable.

Can write this as $F[v] = \bar{F}S^*$, where $\bar{F}[v] = N[v]^{-1}M[v]$ has right parametrix $\bar{G}[v] = M[v]N[v]$:

$N[v] = \mathbf{NMO\ operator}$ $N[v]d(\mathbf{x}_m, h, z) = d(\mathbf{x}_m, h, t(\mathbf{x}_m, h, z))$

$M[v] = \text{multiplication by } A$

$S = \mathbf{stacking\ operator}$

$$Sf(\mathbf{x}_m, z) = \int dh f(\mathbf{x}_m, h, z), \quad S^*r(\mathbf{x}_m, h, z) = r(\mathbf{x}, z)$$

Identify as extension: $\bar{F}[v], \bar{G}[v]$ as above, $X = \{\mathbf{x}_m, z\}, H = \{h\}, \bar{X} = X \times H, \chi = S^*, \eta = S$ - the invertible extension properties are clear.

Standard names for the Standard VA extension objects: $\bar{F}[v]$ = “inverse NMO”, $\bar{G}[v]$ = “NMO” [often the multiplication op $M[v]$ is neglected]; η = “stack”, χ = “spread”

How this is used for velocity analysis: Look for v that makes $\bar{G}[v]d \in \mathcal{R}(\chi)$

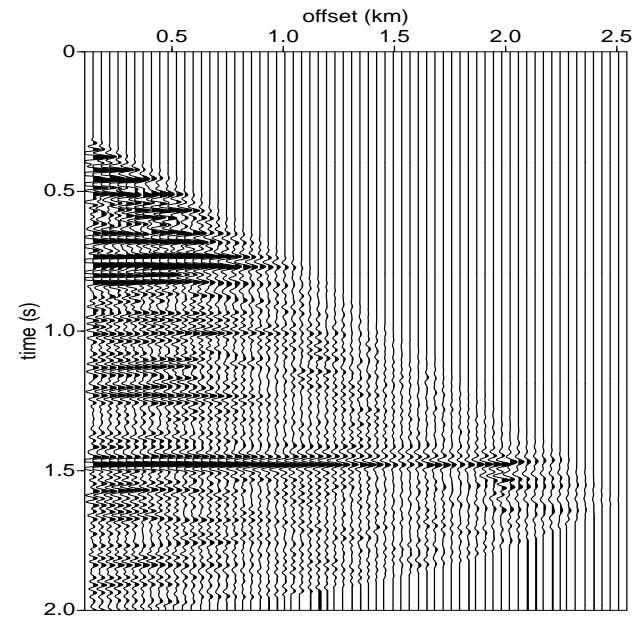
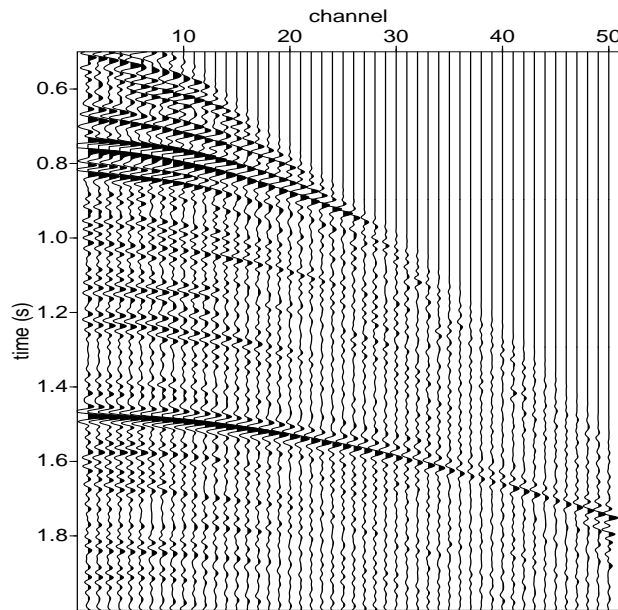
So what is $\mathcal{R}(\chi)$? $\chi[r](\mathbf{x}_m, z, h) = r(\mathbf{x}_m, z)$ Anything in range of χ is *independent of* h . Practical issues \Rightarrow replace “independent of” with “smooth in”.

Inverse problem reduced to: adjust v to make $\bar{G}[v]d^{\text{obs}}$ smooth in h , i.e. *flat* in z, h display for each \mathbf{x}_m (*NMO-corrected CMP*). Same as “feeling for hyperbolas” (Claerbout lectures), layered medium VA in 8.1 of Biondi notes.

Replace z with t_0 , v with v_{RMS} em localizes computation: reflection through $\mathbf{x}_m, t_0, 0$ *flattened* by adjusting $v_{\text{RMS}}(\mathbf{x}_m, t_0) \Rightarrow$ 1D search, do by visual inspection.

Various aids - NMO corrected CMP gathers, velocity spectra, etc.

See Claerbout:BEI, also WWS MGSS 2000 notes for details.



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Left: part of survey (S^{obs}) from North Sea (thanks: Shell Research), lightly preprocessed.

Right: restriction of $\bar{G}[v]S^{\text{obs}}$ to $x_m = \text{const}$ (function of depth, offset): shows rel. sm'ness in h (offset) for properly chosen v .

This only works where Earth is “nearly layered”. Where this fails, go to **Example 2: Surface oriented or standard MVA extension**.

Shot version: Σ_s = set of shot locations, $\bar{X} = X \times \Sigma_s$, $\chi[r](\mathbf{x}, \mathbf{x}_s) = r(\mathbf{x})$.

$$\bar{F}[v]\bar{r}(\mathbf{x}_r, t, \mathbf{x}_s) = \frac{\partial^2}{\partial t^2} \int dx \bar{r}(\mathbf{x}, \mathbf{x}_s) \int ds G(\mathbf{x}_r, t - s; \mathbf{x}) G(\mathbf{x}_s, s; \mathbf{x})$$

Offset version (preferred because it minimizes truncation artifacts): Σ_h = set of half-offsets in data, $\bar{X} = X \times \Sigma_h$, $\chi[r](\mathbf{x}, \mathbf{h}) = r(\mathbf{x})$.

$$\bar{F}[v]\bar{r}(\mathbf{x}_s, t, \mathbf{h}) = \frac{\partial^2}{\partial t^2} \int dx \bar{r}(\mathbf{x}, \mathbf{h}) \int ds G(\mathbf{x}_s + \mathbf{h}, t - s; \mathbf{x}) G(\mathbf{x}_s, s; \mathbf{x})$$

[Parametrize data with source location \mathbf{x}_s , time t , offset \mathbf{h} .] **NB:** note that both versions are “block diagonal” - family of operators (FIOs - tenKroode lectures) parametrized by \mathbf{x}_s or \mathbf{h} .

Properties of surface oriented extension (Beylkin (1985), Rakesh (1988)): if $\|v\|_{C^2(X)}$ “not too big”, then

- \bar{F} has **the Ψ DO property**: $\bar{F}\bar{F}^*$ is Ψ DO
- singularities of $\bar{F}\bar{F}^*d \subset$ singularities of d
- straightforward construction of right parametrix $\bar{G} = \bar{F}^*Q$, $Q = \Psi$ DO, also as generalized Radon Transform - explicitly computable.

Range of χ (offset version): $\bar{r}(\mathbf{x}, \mathbf{h})$ independent of $\mathbf{h} \Rightarrow$ “semblance principle”: find v so that $\bar{G}[v]d^{\text{obs}}$ is independent of \mathbf{h} . Practical limitations \Rightarrow replace “independent of \mathbf{h} ” by “smooth in \mathbf{h} ”.

Application of these ideas = industrial practice of migration velocity analysis.

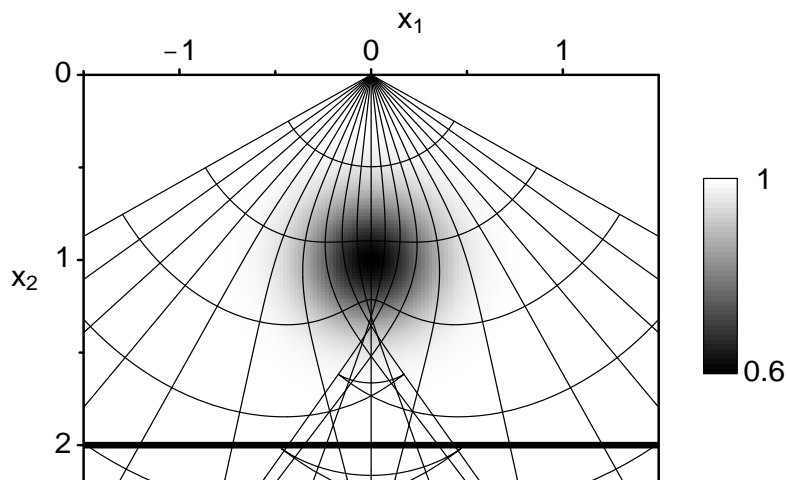
Idea: twiddle v until $\bar{G}[v]d^{\text{obs}}$ is smooth in \mathbf{h} .

Since it is hard to inspect $\bar{G}[v]d^{\text{obs}}(x, y, z, h)$, pull out subset for constant $x, y =$ **common image gather** (“CIG”): display function of z, h for fixed x, y . These play same role as NMO corrected CMP gathers in layered case.

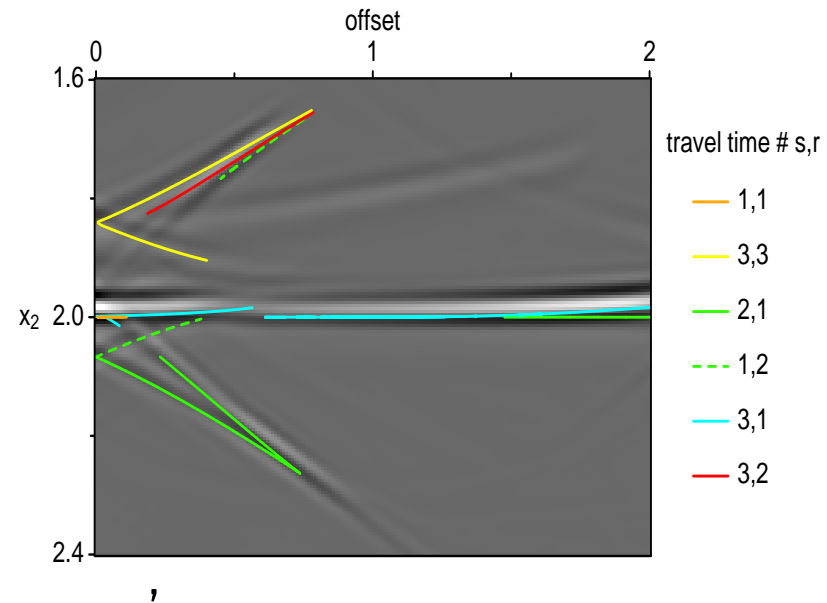
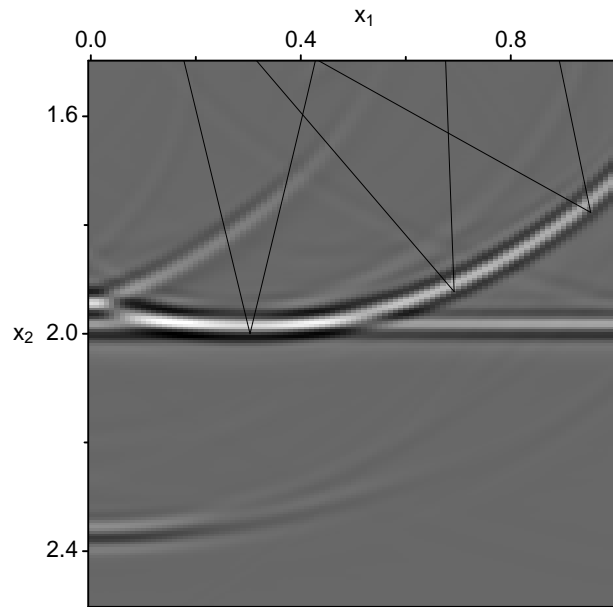
Try to adjust v so that selected CIGs are *flat* - just as in Standard VA. This is much harder, as there is no RMS velocity trick to localize the computation - each CIG depends globally on v .

See Biondi notes for many examples.

Nolan (1997): big trouble! In general, standard extension does **not** have the Ψ DO property. Geometric optics analysis: for $\|v\|_{C^2(X)}$ “large”, multiple rays connect source, receiver to reflecting points in X ; block diagonal structure of $\bar{F}[v] \Rightarrow$ info necessary to distinguish multiple rays is *projected out*.



Example (Stolk & WWS, 2001): Gaussian lens over flat reflector at depth z ($r(\mathbf{x}) = \delta(x_1 - z)$, $x_1 = \text{depth}$).



Left: Const. h slice of $\bar{G}d^{\text{obs}}$: several refl. points corresponding to same singularity in d^{obs} .

Right: CIG (const. x, y slice) of $\bar{G}d^{\text{obs}}$: not smooth in h !

Standard MVA extension only works when Earth has simple ray geometry. When this fails, go to

Example 3: Claerbout's depth oriented extension.

Σ_d = somewhat arbitrary set of vectors near 0 ("offsets"), $\bar{X} = X \times \Sigma_d$, $\chi[r](\mathbf{x}, \mathbf{h}) = r(\mathbf{x})\delta(\mathbf{h})$, $\eta[\bar{r}](\mathbf{x}) = \bar{r}(\mathbf{x}, 0)$

$$\begin{aligned} \bar{F}[v]\bar{r}(\mathbf{x}_s, t, \mathbf{x}_r) &= \frac{\partial^2}{\partial t^2} \int dx \int_{\Sigma_d} dh \bar{r}(\mathbf{x}, \mathbf{h}) \int ds G(\mathbf{x}_s, t-s; \mathbf{x}+2\mathbf{h})G(\mathbf{x}_r, s; \mathbf{x}) \\ &= \frac{\partial^2}{\partial t^2} \int dx \int_{\mathbf{x}+2\Sigma_d} dy \bar{r}(\mathbf{x}, \mathbf{y} - \mathbf{x}) \int ds G(\mathbf{x}_s, t - s; \mathbf{y})G(\mathbf{x}_r, s; \mathbf{x}) \end{aligned}$$

NB: in this formulation, there appears to be too many model parameters.

Computationally economical: for each \mathbf{x}_s solve

$$\bar{F}[v]\bar{r}(\mathbf{x}_r, t; \mathbf{x}_s) = u(\mathbf{x}, t; \mathbf{x}_s)|_{\mathbf{x}=\mathbf{x}_r}$$

where

$$\left(\frac{1}{v(\mathbf{x})^2} \frac{\partial^2}{\partial t^2} - \nabla_{\mathbf{x}}^2 \right) u(\mathbf{x}, t; \mathbf{x}_s) = \int_{\mathbf{x}+2\Sigma_d} dy \bar{r}(\mathbf{x}, \mathbf{y}) G(\mathbf{y}, t; \mathbf{x}_s)$$

$$\left(\frac{1}{v(\mathbf{y})^2} \frac{\partial^2}{\partial t^2} - \nabla_{\mathbf{y}}^2 \right) G(\mathbf{y}, t; \mathbf{x}_s) = \delta(t)\delta(\mathbf{x}_s - \mathbf{y})$$

Finite difference scheme: form RHS for eqn 1, step u forward in t , step G forward in t .

Computing $\bar{G}[v]$: instead of parametrix, be satisfied with adjoint.

Reverse time adjoint computation - specify adjoint field as in standard reverse time prestack migration:

$$\left(\frac{1}{v(\mathbf{x})^2} \frac{\partial^2}{\partial t^2} - \nabla_{\mathbf{x}}^2 \right) w(\mathbf{x}, t; \mathbf{x}_s) = \int dx_r d(\mathbf{x}_r, t; \mathbf{x}_s) \delta(\mathbf{x} - \mathbf{x}_r)$$

with $w(\mathbf{x}, t; \mathbf{x}_s) = 0, t \gg 0$.

Then

$$\bar{F}[v]d(\mathbf{x}, \mathbf{h}) = \int dx_s \int dt G(\mathbf{x} + 2\mathbf{h}, t; \mathbf{x}_s) w(\mathbf{x}, t; \mathbf{x}_s)$$

i.e. exactly the same computation as for reverse time prestack, except that crosscorrelation occurs at an offset $2\mathbf{h}$.

Stolk and deHoop, 2001: Claerbout extension has the Ψ DO property, at least when restricted to \bar{r} of the form $\bar{r}(\mathbf{x}, \mathbf{h}) = R(\mathbf{x}, h_1, h_2)\delta(h_3)$, and under DSR assumption.

Sketch of proof (after Rakesh, 1988):

This will follow from *injectivity* of wavefront or *canonical relation* $C_{\bar{F}} \subset T^*(\bar{X}) - \{\mathbf{0}\} \times T^*(Y) - \{\mathbf{0}\}$ which describes singularity mapping properties of \bar{F} :

$$(\mathbf{x}, \mathbf{h}, \xi, \nu, \mathbf{y}, \eta) \in C_{F_\delta[v]} \Leftrightarrow$$

for some $u \in \mathcal{E}'(\bar{X})$, $(\mathbf{x}, \mathbf{h}, \xi, \nu) \in WF(u)$, and $(\mathbf{y}, \eta) \in WF(\bar{F}u)$

Characterization of $C_{\bar{F}}$:

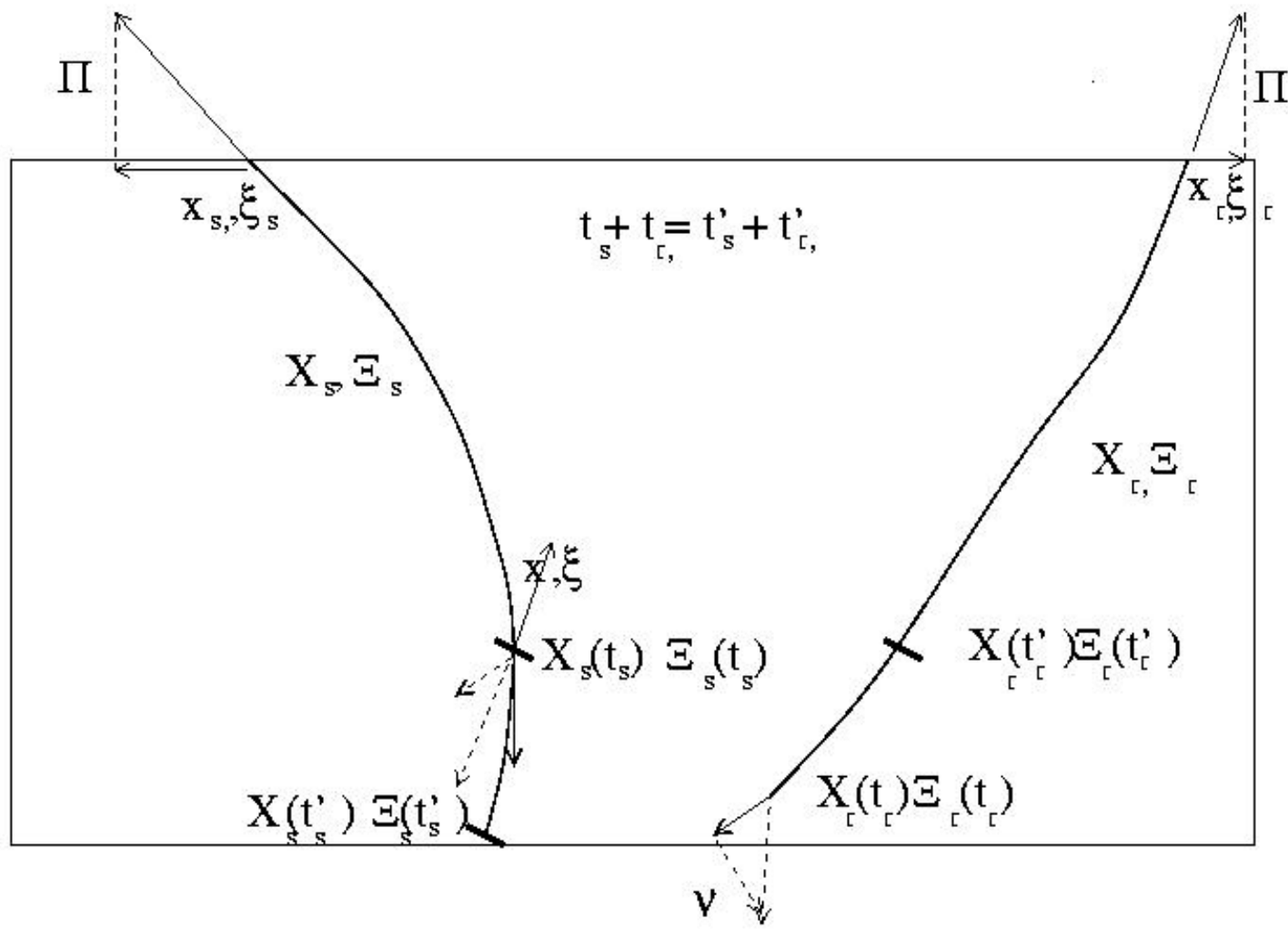
$$((\mathbf{x}, \mathbf{h}, \xi, \nu), (\mathbf{x}_s, t, \mathbf{x}_r, \xi_s, \tau, \xi_r)) \in C_{\bar{F}}[\nu] \subset T^*(\bar{X}) - \{0\} \times T^*(Y) - \{0\}$$

\Leftrightarrow there are *rays of geometric optics* (\mathbf{X}_s, Ξ_s) , (\mathbf{X}_r, Ξ_r) and times t_s, t_r so that

$$\Pi(\mathbf{X}_s(0), t, \mathbf{X}_r(0), \Xi_s(0), \tau, \Xi_r(0)) = (\mathbf{x}_s, t, \mathbf{x}_r, \xi_s, \tau, \xi_r),$$

$$\mathbf{X}_s(t_s) = \mathbf{x}, \mathbf{X}_r(t_r) = \mathbf{x} + 2\mathbf{h}, t_s + t_r = t,$$

$$\Xi_s(t_s) + \Xi_r(t_r) \parallel \xi, \Xi_s(t_s) - \Xi_r(t_r) \parallel \nu$$



Proof: uses wave equations for u, G and

- Gabor calculus: computes wave front sets of products, pull-backs, integrals, etc. See Duistermaat, Ch. 1.
- Propagation of Singularities Theorem

and that's all! [No integral representations, phase functions,...]

Note intrinsic ambiguity: if you have a ray pair, move times t_s, t_r resp. t'_s, t'_r , for which $t_s + t_r = t'_s + t'_r = t$ then you can construct two points $(\mathbf{x}, \mathbf{h}, \xi, \nu), (\mathbf{x}', \mathbf{h}', \xi', \nu')$ which are candidates for membership in $WF(\bar{r})$ and which satisfy the above relations with the same point in the cotangent bundle of $T^*(Y)$.

No wonder - there are too many model parameters!

Stolk and deHoop fix this ambiguity by imposing two constraints:

- DSR assumption: all rays carrying significant reflected energy (source or receiver) are upcoming.
- Restrict \bar{F} to the domain $\mathcal{Z} \subset \mathcal{E}'(\bar{X})$

$$\bar{r} \in \mathcal{Z} \Leftrightarrow \bar{r}(\mathbf{x}, \mathbf{h}) = R(\mathbf{x}, h_1, h_2)\delta(h_3)$$

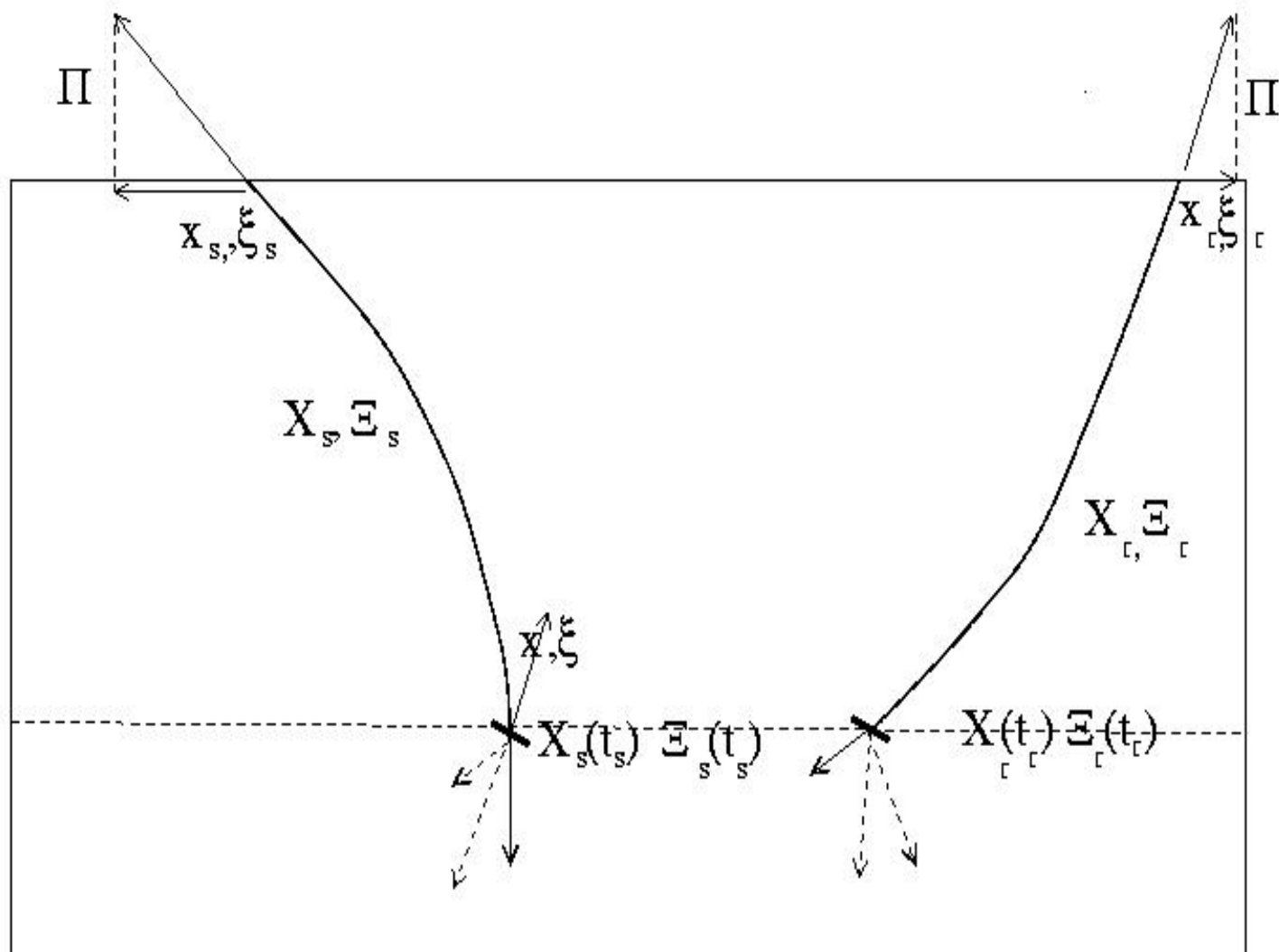
If $\bar{r} \in \mathcal{Z}$, then $(\mathbf{x}, \mathbf{h}, \xi, \nu) \in WF(\bar{r}) \Rightarrow h_3 = 0$. So source and receiver rays in $C_{\bar{F}}$ must terminate at same depth, to hit such a point.

Because of DSR assumption, this fixes the traveltimes t_s, t_r .

Restricted to \mathcal{Z} , $C_{\bar{F}}$ is injective.

$$\Rightarrow C_{\bar{F}^* \bar{F}} = I$$

$\Rightarrow \bar{F}^* \bar{F}$ is Ψ DO when restricted to \mathcal{Z} .



Quantifying the semblance principle: devise operator W for which

$$\ker W \simeq \mathcal{R}_\chi,$$

then minimize a suitable norm of

$$W\bar{G}d^{\text{obs}}.$$

Converts inverse problem to optimization problem. With proper choice of W , Ψ DO property \Rightarrow objective is *smooth* \Rightarrow can use Newton and relatives.

Upshot: Claerbout's depth oriented extension appears to offer basis for efficient new algorithms to solve velocity analysis problem - research currently under way in several groups.

Summary:

- quite a bit is known about the imaging problem under “standard hypotheses”: mathematics of multipathing imaging (asymptotic inversion, invertible extensions) clarified over last 10 years.
- many imaging situations (eg. near salt - cf. Biondi) violate “standard hypotheses” grossly - need much better theory
- extension of imaging *via* multiple suppression - some progress, many open questions re non-surface multiples
- velocity analysis - some progress, but still in primitive state mathematically
- almost no progress on underlying nonlinear inverse problem